The ur waters. л пээмцэр lysimeters water. Wa ewied gai qsmbenec itisoqmoo Although 3 cont Georgia) ter. D. We to gaixim apparent flow duri cated tha nuott pr tion over nob-bnsi water mi: Universit V.l.woff cating ch water ch piqotosi и цапоц n eviensi

among flowpath, isotopic signature, travel tuff. The work focused on the relations field site in Arizona in fractured unsaturated chemical indicators of flow pathways at a oral session with a look at isotopic and associates (University of Arizona) ended the groundwater (high 222Rn). R. Bassett and saturated zone water (low 222Rn) and soil 222 Rn was used to distinguish between unon the basis of its high Ca content, and tinguished from the two soil end-members In this case, bedrock groundwater was disgeneration over a wide variety of conditions. mixing model is appropriate for streamflow Ca suggest that a simple three-component, surements of naturally occurring 222Rn and Walker Branch Watershed, Tenn. Their meamodel for streamflow generation at the Ca as tracers in a three-component mixinginformation concerning the use of  $^{222}\mathrm{Rn}$  and chusetts Institute of Technology) presented sheds. D. Genereux and associates (Massaof carbonic acid weathering in the waterprimarily reflect differences in the amounts

known and the gas transfer rate from the that the concentration in groundwater is tions of streamflow components, provided that <sup>222</sup>Rn is useful for mass-balance separaterrain in central Florida. They demonstrated flow losses along stream reaches in karst to determine groundwater gains and streamstrating the use of 222 Rn and introduced 5F<sub>6</sub> versity of Florida) presented work demonend-member. K. Ellins and associates (Unigroundwater had been used as the "old" be calculated if base flow rather than 95% of peak flow); different amounts would contained large amounts of old water (55natures. He showed that meltwater events showed large spatial variability in 818 sig-N.Y., watersheds. Both soil and groundwater sources of peak flow in two Adirondack, composition of stream water to determine Burns (USGS) examined the use of 8<sup>18</sup>O teresting summaries of isotopic research. D. Poster presentations included several intunnel of a copper mine. flow systems of fluid entering an ore haulage

vealed the presence of two interconnected

time, fracture recharge, perched water zones,

and surface drainage. Isotope tracing re-

vited paper and summarized the five major mechanisms for rapid delivery of pre-event water to stream channels, including translatory flow, macropore flow, groundwater ridging, extension of saturated conditions in more conductive near-surface layers, and the roles of these different mechanisms in different environments, Buttle advocated more complete integration of environmental isotope studies with the appropriate hydrometric measurements than has been achieved previously.

L. Cooper (Oak Ridge National Laborawere not sustainable for their system. of displaced old water relative to new water conventional explanations for large amounts soil. This result led them to assume that ilar to the background concentration in the δι8Ο composition of drainage water was simwater controversy. They showed that the highlighting implications for the old-/newdrains on a drained grassland soil in Britain, sults of an isotope tracer experiment in mole (Polytechnic Southwest, U.K.) presented retial soil water. P. Addison and associates matrix waters, which were similar to the inito the applied concentrations, differed from concentrations of  $\delta^{18}$ O and Br, while similar samples and matrix water. Lysimeter tracer comparing the compositions of lysimeter adding tracers to intact soil cores and then representativeness of lysimeter samples by west, U.K.) and associates investigated the tions to flow. J. Dowd (Polytechnic Southindicated much larger old-water contribu-Si and Cl separations, although similar, both new and old water determined with 818O, the slow flow closely matched the amounts of though the relative amounts of quick and values indicated was mainly new water. Alall the flow was quickflow, which the 818O catchment in China and found that almost using 8180, Si, and Cl at a small artificial model with hydrograph separations made associates compared a unit hydrograph C. Kendall (U.S. Geological Survey) and

ing to streamflow. Interstream differences trom deep and shallow flowpaths contributing rapid changes in the amounts of water trastorm oscillations in composition, indicaters showed rapid and highly correlated instreams at Catoctin Mountain, Md. Both tracand 813C as tracers of water flowpaths at two Kendall (USGS) examined the use of \$75r/85r with evaporative effects. T. Bullen and C. preferential melt and refreezing, combined fractionation included a complex pattern of tractionation during snowmelt. The apparent but it was also consistent with significant was indicated by physical measurements, of melt waters from snow and soil ice than isotope work showed a much greater mixing and streamtiow in an Arctic watershed. Their physical measurements to monitor snowmelt [0012][000c][0004]the use of 5<sup>18</sup>O and tory) and associates combined

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HADROFOGA

SECLION NEMS

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## Isotope Tracers in Hydrology

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A major uncertainty in hydrologic and chemical modeling of watersheds has been the quantification of the contributions of watersheds has been ter and solutes from various hydrologic pathershould models exist that can reproduce the timing and quantity of the produce the timing and quantity of the coupled with these physical models, they do not accurately reflect the streamwater chemistry. Therefore, it has been concluded that istry. Therefore, it has been concluded that either the hydrologic pathways are represented inaccurately in the models or the processes affecting the water chemistry are not well understood.

Well understood.

To provide a forum for a critical evaluarically portray catchment hydrology. whether simple mass-balance models realisof water components are negligible, or and temporal variation in the compositions behave conservatively, whether the spatial stormflow, whether isotope tracers actually contributions of "old water" to channel sonuqa the processes responsible for rapid ertheless, considerable debate still sursome mix of event and pre-event water. Nevin the catchment prior to the event, but with snowmelt is supplied largely by water stored streamflow generated during rainfall or ing models have shown repeatedly that by simple two-component, conservative-mixlsotope hydrograph separations determined climatic and hydrogeologic environments. quess these concerns in a wide variety of

tion of the use of isotope tracers in catching a special session of ACU's ment hydrology, a special session of ACU's 1991 Fall Meeting in San Francisco, "Evaluation of Streamflow Generation Using Isotope Techniques," was convened. It was organized and co-chaired by Jeff McDonnell and Carol Kendall.

Carol Kendall.

Two invited papers opened the oral session invited papers.

sion. M. Sklash (University of Windsor) provided a thorough review of environmental isotope studies of runoff generation and the simple mass-balance approach. He then outlined some difficulties associated with the simple two-component approach and new sensitivity analyses to determine the effectiveness of appropriate approaches. J. Buttle tiveness of appropriate approaches. J. Buttle (Trent University) presented the second in-

shallow soil horizons. K. Rice and O. Bricker

old-water flux through highly conductive

stream, due in part to underestimation of

of groundwater flux could not account for

groundwater sources. Physical measurement

ter, they advocated the use of a three-com-

Si concentration were observed in groundwa-

(40-85%). Because large spatial variations in

water contributions to peak flow were high

pic, chemical, and physical measurements. Use of both  $\delta^{18}$ O and Si showed that old-

a headwater stream in Ontario using isoto-

100) examined groundwater contributions to

Hinton and associates (University of Water-

correct for loss of 222 Rn from the stream. M.

stream to the atmosphere is measured to

their observed old-water volumes in the

ponent model with shallow and deep

mounts vater-(Massaesented <sup>22</sup>Rn and mixingthe neir mea-Rn and ponent eamflow inditions. vas disembers and een unnd soil tt and nded the and ys at a saturated ons :, travel ater zones, ng rennected re haulage

everal insearch, D. δ<sup>18</sup>Ο termine ıdack, oundwater δ18O sigr events ater (55nts would ıan : "old" ites (Unidemonduced SF<sub>6</sub> nd streamin karst nonstrated ce separaprovided ater is om the ured to stream. M. of Waterbutions to ng isotoirements. hat oldvere high ariations in groundwaree-comeep asurement ount for in the ation of ductive 1 O. Bricker ta from I Maryland. an integra-

cating ex-

tensive mixing in the unsaturated zone. Although minor spatial variation in soil-water isotopic composition was observed, soilwater chemistry was spatially variable, indicating chemical reactions during unsaturated flow. J. Waddington and N. Roulet (York University) examined groundwater-surface water mixing during storms for a small wetland-dominated watershed. Although saturation overland flow appeared to dominate nunoff production, δ<sup>18</sup>O and LiBr tracers indicated that prestorm water dominated stormflow during artificial sprinkling events. The apparent dichotomy was resolved by rapid mixing of rainwater and emerging groundwater. D. Wenner and associates (University of Georgia) investigated the changes in the δ<sup>18</sup>O content of recharge waters in Georgia. Although the rain was variable in isotopic composition, this variability was significantly dampened in shallow soil waters due to mixing between new rain and older immobile water. Waters collected from zero-tension lysimeters were intermediate in composition between matrix water and tension lysimeter

The underlying theme of most of the pre-

sentations was that the isotopic compositions of rain, throughfall, soil water, and groundwater are commonly variable in time and space. The general consensus was that if such variability is significant at the catchment scale, then simple two- or three-component, constant-composition end-member mixing models (that is, the "classic" isotope hydrograph-separation technique) may not provide realistic interpretations of the system hydrology. The various presentations provided novel ways of confirming or refining hydrologic models derived largely from oxygen and hydrogen isotope data by combining these data with other approaches, including applying tracers to the system, integrating isotope hydrograph models with appropriate hydrometric measurements, trving other isotope tracers (strontium, carbon, and radon isotopes), combining chemical and isotopic data, and using multicomponent, end-member mixing models.

A panel discussion at the end of the morning session raised several questions regarding the use of isotope tracers in hydrology. A major concern was how to separate old- and new-water components or gain information on the relative movement of wa-

ter along various flowpaths when the isotopic compositions of the water sources are so often variable in composition. Several presenters and audience participants are actively trying to demonstrate the nature of this variability at different spatial and temporal scales. Several unanswered questions were raised, including: Are multicomponent oldwater models warranted; Why do δ<sup>18</sup>O, Cl, and Si separations often agree, despite the obvious lack of conservativeness displayed by Cl and Si in shallow and deep systems; Is there still a rationale for studies making the assumption that rainfall has a constant isotopic composition; and Have hydrologists ignored important evaporative effects in snow, soil water, and throughfall in small watershed tracing studies? We hope that many of the remaining questions will be addressed at the "Tracers in Hydrology" session at the International Association of Hydrological Sciences meeting to be convened July 11-23, 1993, in Yokohama, Japan.—Jeffrey J. McDonnell, Utah State University, Logan; and Carol Kendall, U.S. Geological Survey, Menlo Park, Calif.